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EXAMPLES OF THE MEAN THREE-DIMENSIONAL WIND CIRCULATION OVER THE TROPICAL NORTH PACIFIC OCEAN

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ABSTRACT

Two extreme climatological contrasts in the 1954 mean monthly wind circulation over the tropical North Pacific Ocean are presented in the form of cross sections and horizontal analyses. The wind field is shown on cross sections near 165° E. from 5° to 35° N. latitude. The horizontal variations of the wind are shown at 10,000-foot intervals throughout the troposphere and lower stratosphere. Twenty-seven figures are presented.

1. INTRODUCTION

Until recently, the lack of upper air data in the Tropics has prevented climatologists from preparing maps depicting mean high-level circulation patterns over most of the tropical North Pacific Ocean. In an attempt to correct the lack, as far as this is possible with recent RAWIN data, one phase of the work conducted in tropical meteorology at the Oahu Research Center has been the drawing of detailed analyses of mean monthly winds for various atmospheric levels over the Marshall Islands (Korshover [1]). Although limited data have confined Korshover's work to the atomic weapons testing periods consisting of GREENHOUSE (spring 1951), IVY (fall 1952), and CASTLE (spring 1954), study of the analyses has suggested that two, significantly different, mean wind flow patterns regularly occur over the Marshall Islands. Consequently, it is the purpose of this paper to describe these patterns more fully by extending the three-dimensional analysis of the mean wind flow patterns to the entire tropical North Pacific Ocean area.

2. THE TWO MEAN WIND FLOW PATTERNS

The first mean wind flow pattern illustrated (figs. 8-17) is derived from winds prevailing during the month of February 1954. This was a "high index" period and the horizontal analyses are believed to be typical at each level of the mean wind circulation over the tropical North Pacific Ocean area during high zonal index in temperate latitudes. Daily analyses show that the synoptic flow pattern in the middle and high troposphere through the month was essentially zonal and strong in speed.

The second mean wind flow pattern illustrated (figs. 18-27) is based on winds observed during the period April 16 through May 15, 1954. For this interval of time the mean horizontal analyses are believed to be typical at each level studied of the mean wind circulation during a "low index" period. That the mean wind analyses are representative of high and low index situations can be checked by the striking similarity of the climatological maps to the daily synoptic maps for the same periods as analyzed by the weather personnel on OPERATION CASTLE. In other words, the daily departure of the synoptic field of motion, from the mean, for both periods was very small.

The meager evidence from years before 1954 (almost

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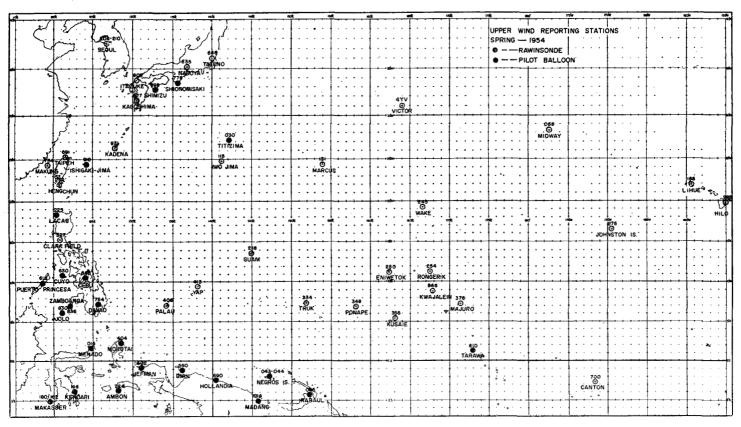


FIGURE 1.—Upper-air observing network in the Central Pacific during spring 1954.

Table 1.—Pacific upper-wind stations, types of wind finding equipment, and times of observation

| Station No. | Station name | Type of wind finding equipment** | Time of observa- | Station No. | Station name | Type of wind finding equipment** | Time of observa- | | | | | |
|--|--|----------------------------------|--|---|--|----------------------------------|--|--|--|--|--|--|
| | Block 91 | | | Block 47 | | | | | | | | |
| 066. 115. 131. 4 Y V 165. 218. 245. 2250* 254. 275. 285. 334. 334. 3366* 366* 376* 408. 413. 530. 533. 610. 623. 487. 490. 700. 705. | Midway Iwo Jima Marcus. Victor. Lihue. Guam Wake. Eniwetok Rongerik Johnston. Hilo. Truk Ponape Kusale Kwajalein Majuro Palau Yap. Nauru Ocean. Tarawa Beru Fanning Christmas Canton Gardner. Hull | REERERERERERERERECOPOCORCO | 03, 09, 15, 21 03, 09, 15, 21 00, 06, 12, 18 00, 06, 18 00, 06, 18 00, 06, 18 00, 06, 18 00, 06, 18 | 208. 635. 646. 778. 808. 827. 898. 931. 014. 085. 014. 146. 180. 404. 502. 560. 690. 724. | Seoul Nagoya Tateno Shionomisaki Itazuke Kagoshima Shimizu Okinawa Block 94 Madang Rabaul Block 97 Menado Kendari Makasser Moritai Jefman Biak Hollandia Ambon Block 98 | RRRRRRRR | 03, 09, 15, 21 03, 09, 15, 21 04, 10, 16, 22 04, 10, 16, 22 04, 10, 16, 22 05, 11, 23 05, 11, 23 05, 11, 23 02, 08, 23 03, 08, 22 02, 08, 22 05, 11, 23 | | | | | |
| 10 | Block 46 | OL. | 00,00,18 | 223 327 618 630 | Laoag Clark Field Puerto Princesa Cuyo | P R P | 03, 15 03, 09, 15, 21 03, 15, 21 03, 21 | | | | | |
| 697 734 752 | Taipeh Makung Hengchun | R R R | 03, 09, 15, 21 03, 09, 15, 21 03, 09, 15, 21 | 645 | Cebu | P P P P | 03, 15 03, 15 03, 15 03, 09 03, 15, 21 | | | | | |

 $[\]begin{tabular}{ll} \begin{tabular}{ll} \beg$

Table 2.—Number of wind observations at each wind analysis level

| | Level (ft.) | | | | | | | | | | Level (ft.) | | | | | | | | | - 11-1-1 | |
|--|--|---|---|---|---|---|--|--|--|-------------------------------|-------------|---|---|---|---|--|---|--|--|--|--|
| Station | 2,000 | 10,000 | 20,000 | 30,000 | 40,000 | 50,000 | 60,000 | 65,000 | 70,000 | 80,000 | Station | 2,000 | 10,000 | 20,000 | 30,000 | 40,000 | 50,000 | 60,000 | 65,000 | 70,000 | 80,000 |
| FEBRUARY 1-28, 1954 Block 91 | | | | | | | | | | | API | RIL 16- Blo | MAY | 15, 1954 | | <u> </u> | | 1 | | | |
| 4YV | 80 108 90 106 97 111 108 123 82 112 90 54 106 101 117 107 52 47 | 79 110 69 106 74 111 110 123 87 110 47 106 108 109 42 41 23 | 50 106 53 107 64 108 107 122 85 109 61 30 108 107 81 110 23 32 | 17 88 41 91 47 106 98 123 85 108 109 71 108 10 26 | 644 288 711 344 105 79 123 85 100 37 26 108 109 66 107 7 26 | 41 18 38 23 89 44 122 79 89 27 22 107 58 104 | 24 10 18 10 49 17 94 56 58 16 71 75 48 93 | 20 9 8 10 45 13 92 51 49 15 69 64 41 91 15 | 17 10 39 86 51 45 16 69 63 38 89 | 111 9 26 64 42 37 63 57 29 85 | 4YV | 103 118 94 115 114 115 146 117 110 57 130 126 139 125 58 57 73 120 | 95 111 104 114 78 114 117 83 54 130 129 128 50 43 32 114 | 84 106 104 111 67 114 116 142 117 71 130 128 41 30 | 64 102 104 103 63 110 116 140 117 60 22 130 129 137 127 30 25 | 46 85 104 99 60 105 111 140 114 53 20 129 129 126 26 24 | 32 50 96 86 33 98 95 137 72 36 18 129 113 103 121 23 21 | 21 27 62 65 19 63 66 109 39 24 13 108 84 58 19 | 22 57 59 17 58 51 99 30 14 97 76 51 88 18 | 177 500 50 15 53 43 891 277 12 94 74 48 86 | 40 35 35 16 81 26 89 72 34 77 |
| 700 112 105 77 63 62 49 26 21 13 Block 46 | | | | | | | | Block 46 | | | | | | | | | | | | | |
| 697 734 | 59 73 73 | 53 65 71 | 47 59 64 | 38 37 61 | 22 62 | 41 | | | | · | 697 752 | 49 79 | 78 | 76 | 78 | 75 | 67 | | | | |
| 782 | | | | | | | | | Block 47 | | | | | | | | | | | | |
| 208 | 69 101 47 100 94 91 89 86 | 71 82 48 90 60 64 62 55 | 64 61 48 72 51 57 22 39 | 56 49 44 40 42 19 | 41 29 34 26 | 29 21 19 | | | | | 208 | 89 102 55 99 106 92 80 108 | 87 90 56 83 101 75 55 107 | 89 72 54 58 95 55 | 81 53 52 46 67 27 | 56 43 41 24 46 | 25 26 23 64 | | | | |
| Block 94 | | | | | | | Block 94 | | | | | | | | | | | | | | |
| 014 044 085 | 43 12 38 | 31 | | | | | | | | | 014 | 39 46 | 23 25 | | | | | | | | |
| Block 97 | | | | | | | | Block 97 | | | | | | | | | | | | | |
| 014 | 38 44 44 21 20 44 27 55 | 23 19 | | | | | | | | | 014 | 42 24 53 26 42 37 65 37 | 26 18 34 22 29 20 21 21 | | | | | | | | |
| Block 98 | | | | | | | Block 98 | | | | | | | | | | | | | | |
| 223 | 41 88 33 36 47 47 47 47 | 40 76 26 30 34 34 33 39 | 24 60 23 22 | 16 51 | 48 | 38 | | | | | 223 | 49 114 25 47 48 43 18 41 | 46 105 18 30 42 31 14 39 | 31 83 33 19 29 | 18 62 | 59 | 57 | | | | |

entirely confined to periods during which tests were made in the Marshalls) strongly suggests that, if a long series of observations were available, the two situations presented here as "high" and "low" index would be typical in the tropical Pacific of the months February and May; that is, that anticyclogenesis aloft in the Central Pacific culminates in February, cyclogenesis in May. However, the voluminous data upon which a firm statement about seasonal trends could be made have yet to be collected. For this reason, the terms "high" and "low" index, as used here, are purely qualitative and descriptive, and no attempt is made to disentangle the synoptic from the seasonal aspects of the two situations. In the high-index situation, the jet stream is strong in speed, the subtropical anticyclones aloft lie close to the equator, and the entire flow pattern is strongly zonal. In the low-index situation the jet stream is weak, the subtropical anticyclones aloft are centered outside

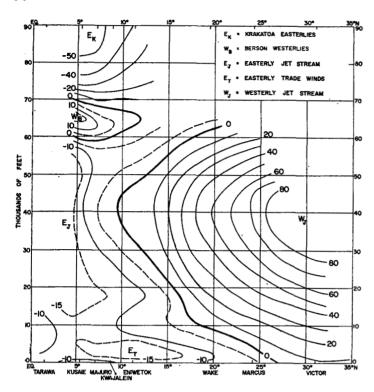


Figure 2.—Mean zonal component of the wind in knots near 167° E., February 1954.

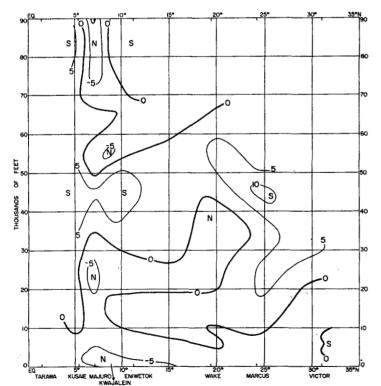


FIGURE 3.—Mean meridional component of the wind in knots near 167° E., February 1954.

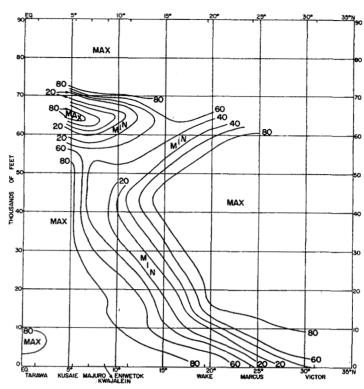


FIGURE 4.—Steadiness of the wind in percent near 167° E., February 1954.

the Tropics, and the wind pattern aloft shows large meridional air movement associated with upper-level cyclogenesis. The maps presented here illustrate the low-latitude concomitants of these situations. The difficult question of the part played by seasonal variation has to be treated at greater length in future publications.

3. ANALYSIS METHOD

The area of analysis extended from Hawaii to the coast of Asia in the east-west direction and from 0° to 35° N. in the north-south direction. The lowest level of analysis selected was 2,000 feet; winds at that level are not affected by possible orographic disturbances due to low-lying atolls. Good evidence of profound orographic effects is seen near the borders of the maps, such as in the Philippines, Formosa, and Japan. Here, the 2,000- and even 10,000-foot winds reveal large departures from a consistently smooth flow.

Above 2,000 feet, and starting with the 10,000-foot chart, the wind analyses have been carried out at 10,000-foot intervals to 80,000 feet (figs. 8-27). Because the wind structure immediately above the tropopause is complex, an additional chart at 65,000 feet has been included.

All upper wind data were obtained from teletypewriter collectives received at various stations in the Pacific. The upper wind reporting stations, their type of wind-finding equipment, and their scheduled times of observa-

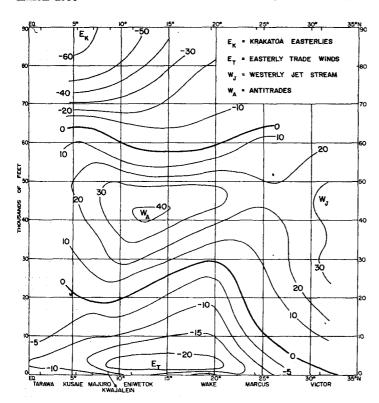


FIGURE 5.—Mean zonal component of the wind in knots near 167° E., April 16-May 15, 1954.

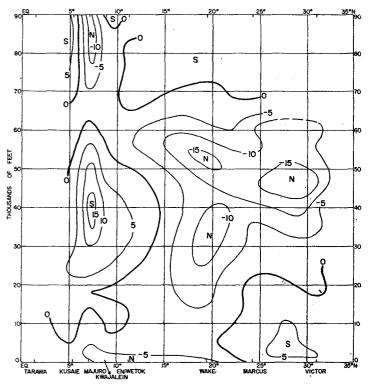


Figure 6.—Mean meridional component of the wind in knots near 167° E., April 16-May 15, 1954.

tion are listed in table 1 for reference. The geographic distribution of the stations is shown in figure 1. Table 2 contains the number of wind observations that were used to arrive at the mean resultant wind value for each analysis level.

The streamline analyses have been performed by the isogon method (Palmer and others [2]). The speed lines have been drawn for 10-knot intervals, and occasionally where gradients were weak, a 5-knot interval was used.

The three-dimensional aspect of the wind circulation shown on each horizontal wind chart may be enhanced through the use of a wind cross-section drawn for 167° E. The \overline{u} and \overline{v} components of the wind and the wind steadiness $(\overline{S} = \frac{|\overline{\mathbf{V}}|}{\overline{C}})$ where S is the steadiness, $|\overline{\mathbf{V}}|$ is the

mean vector speed, and \overline{C} is the mean wind speed irrepective of wind direction) are presented in figures 2-7. The names of all large-scale wind systems have been identified at their centers with appropriate symbols (Palmer [2], [3]).

4. COMMENTS ON THE MAPS

The following points deserve special emphasis:

- 1. No other tropical region in the world contains so dense a RAWIN observational network as that now in existence in the Central Pacific.
 - 2. The wind circulations presented herein were selected

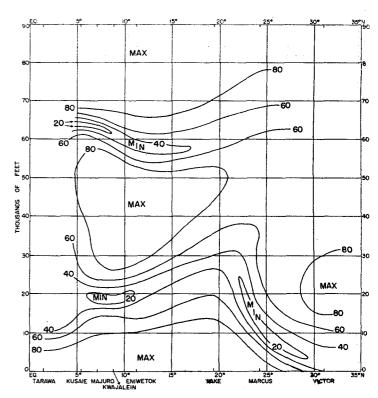


FIGURE 7.—Steadiness of the wind in percent near 167° E., April 16-May 15, 1954.

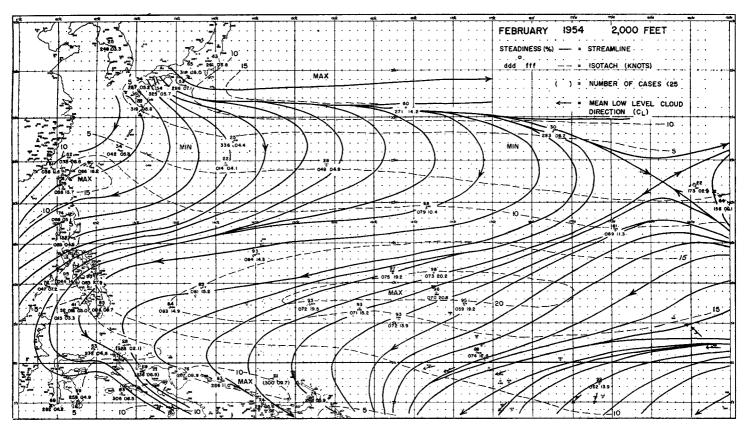


FIGURE 8.—The field of motion at 2,000 feet, February 1954.

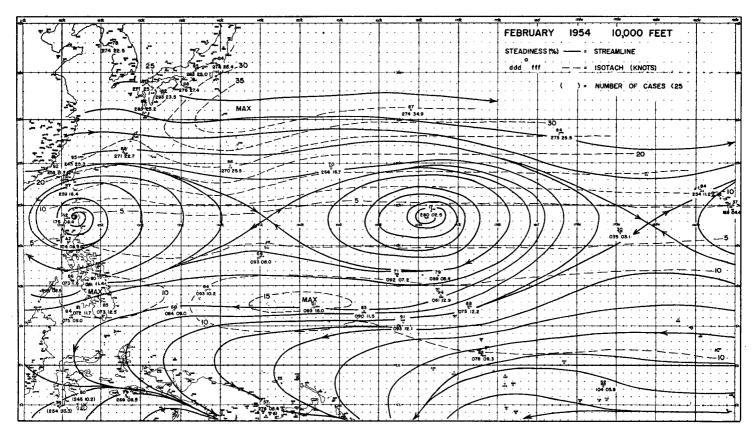


FIGURE 9.—The field of motion at 10,000 feet, February 1954.

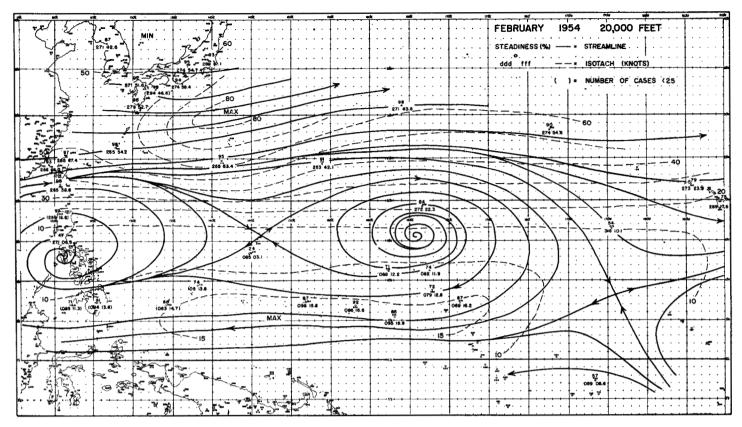


FIGURE 10.—The field of motion at 20,000 feet, February 1954.

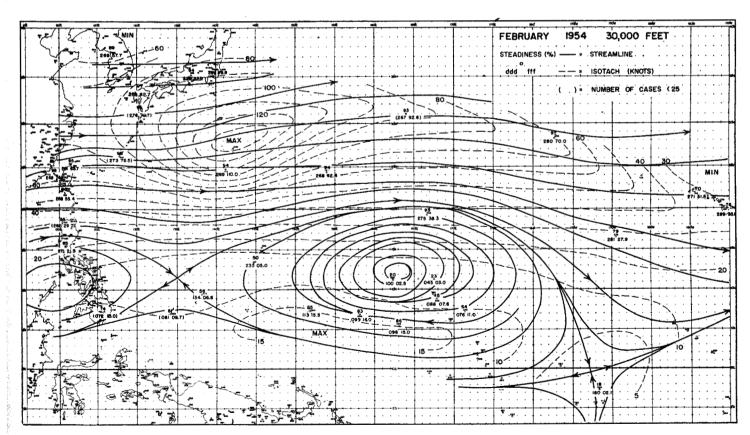


FIGURE 11.—The field of motion at 30,000 feet, February 1954.

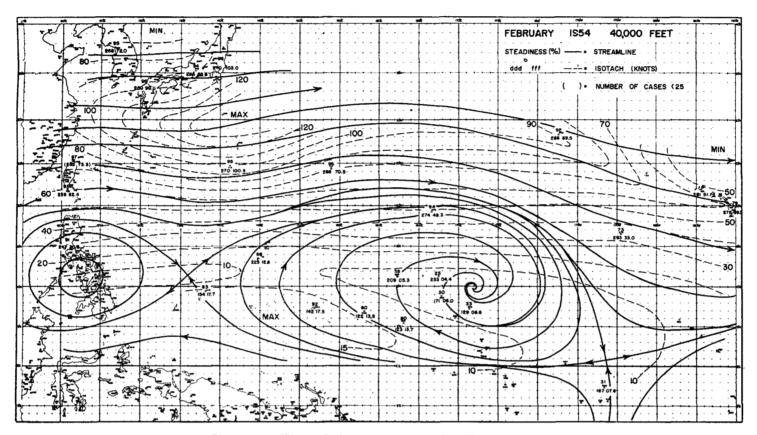


FIGURE 12.—The field of motion at 40,000 feet, February 1954.

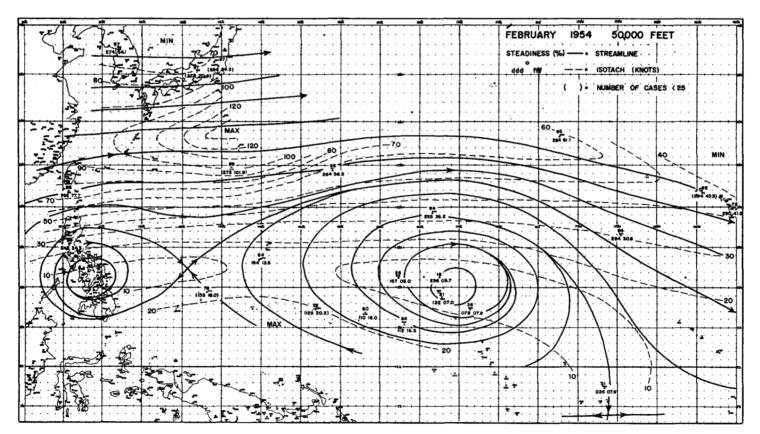


FIGURE 13.—The field of motion at 50,000 feet, February 1954.

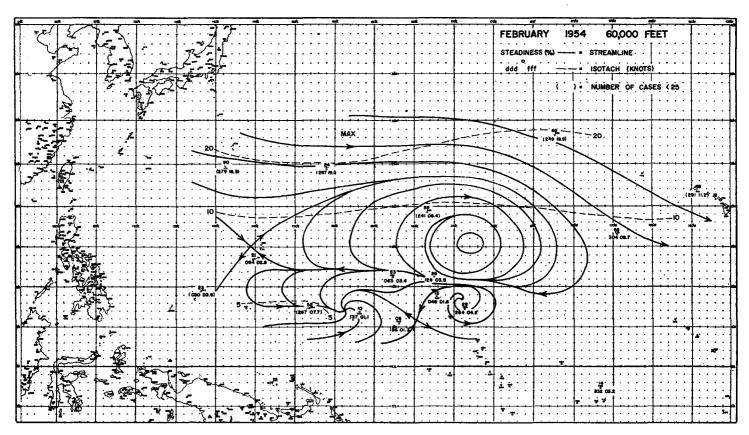


FIGURE 14.—The field of motion at 60,000 feet, February 1954.

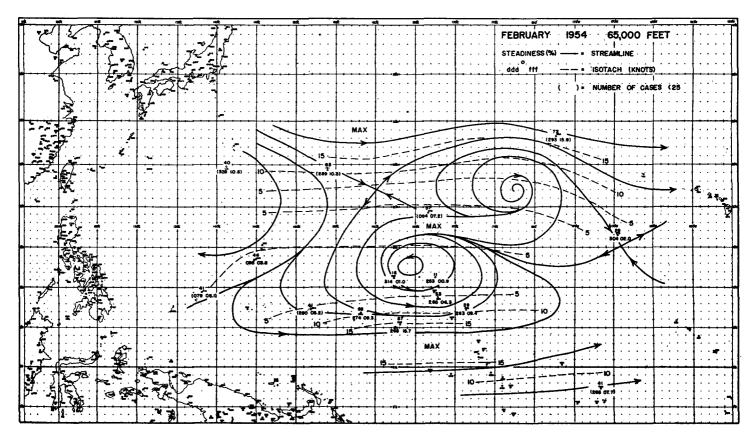


FIGURE 15.—The field of motion at 65,000 feet, February 1954.

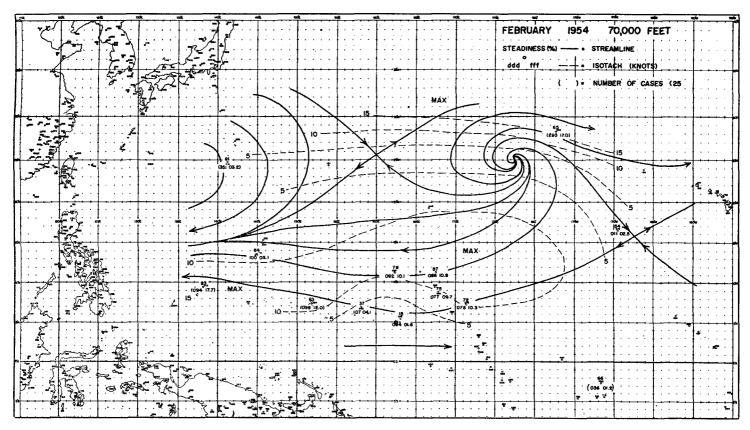


FIGURE 16.—The field of motion at 70,000 feet, February 1954.

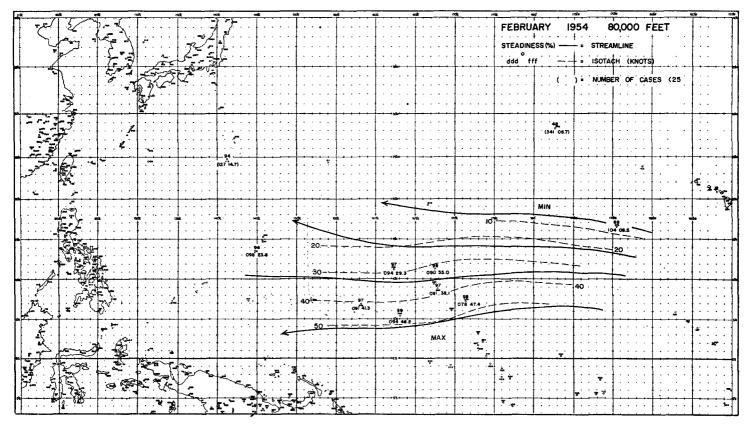


Figure 17.—The field of motion at 80,000 feet, February 1954.

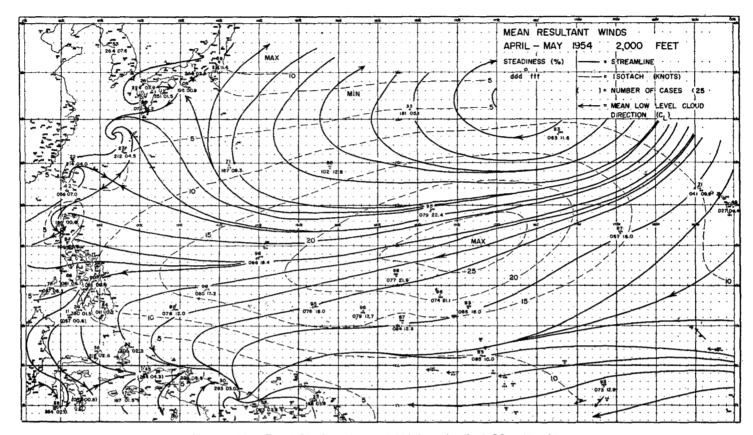


FIGURE 18.—The field of motion at 2,000 feet, April 16-May 15, 1954.

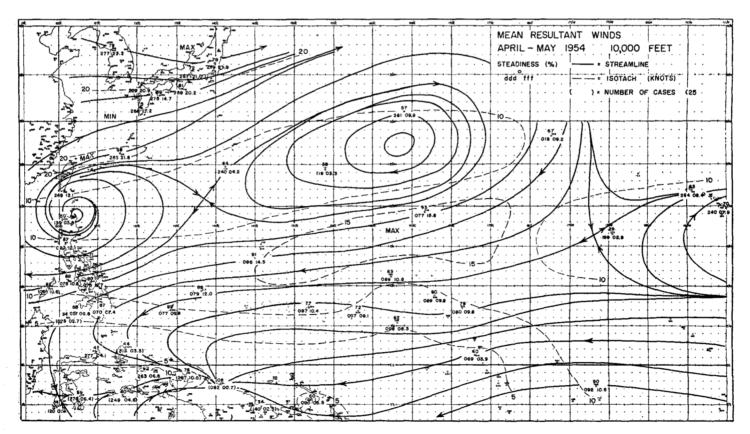


FIGURE 19.—The field of motion at 10,000 feet, April 16-May 15, 1954.

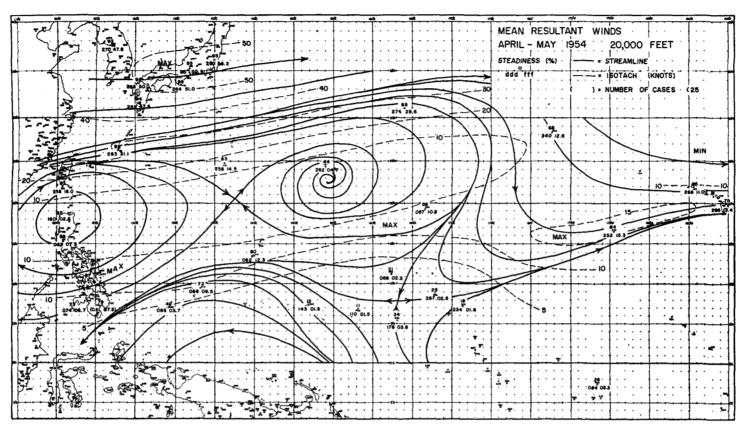


FIGURE 20.—The field of motion at 20,000 feet, April 16-May 15, 1954.

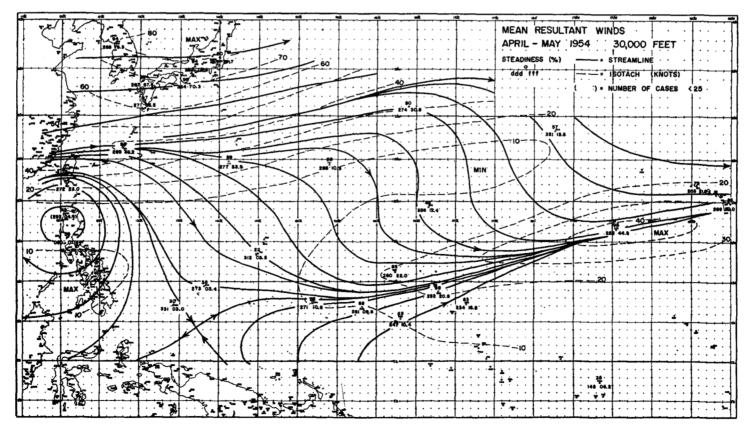


FIGURE 21.—The field of motion at 30,000 feet, April 16-May 15, 1954.

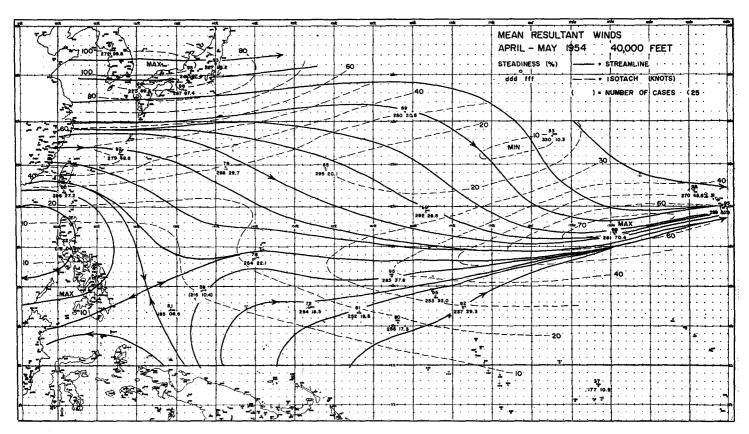


FIGURE 22.—The field of motion at 40,000 feet, April 16-May 15, 1954.

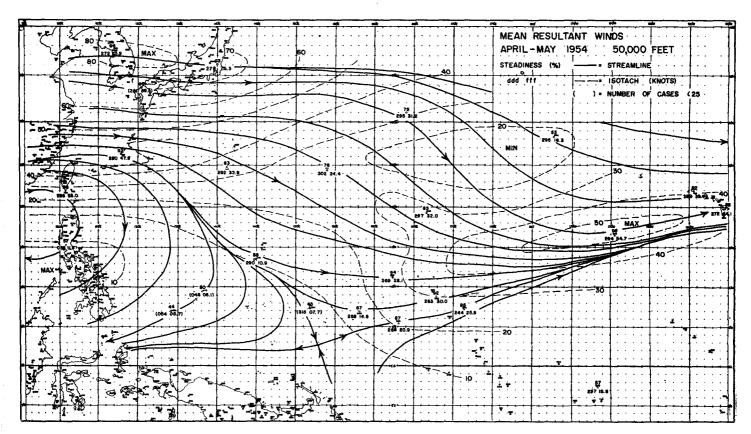


FIGURE 23.—The field of motion at 50,000 feet, April 16-May 15, 1954.

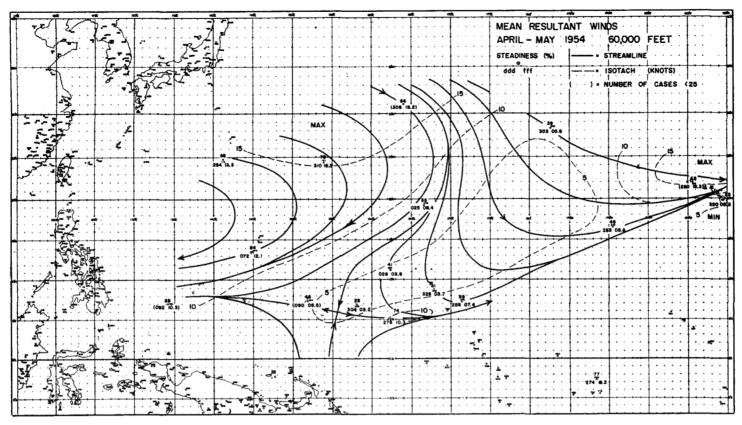


FIGURE 24.—The field of motion at 60,000 feet, April 16-May 15, 1954.

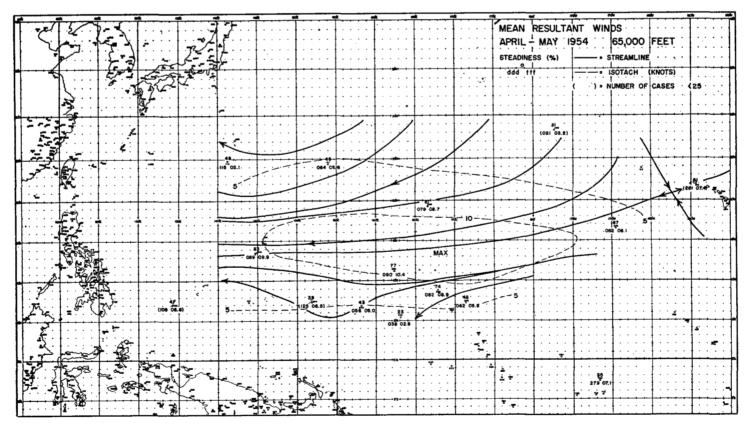


FIGURE 25.—The field of motion at 65,000 feet, April 16-May 15, 1954.

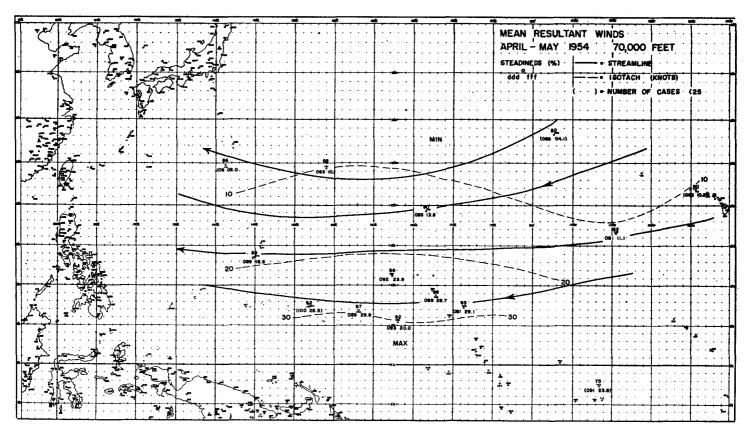


FIGURE 26.—The field of motion at 70,000 feet, April 16-May 15, 1954.

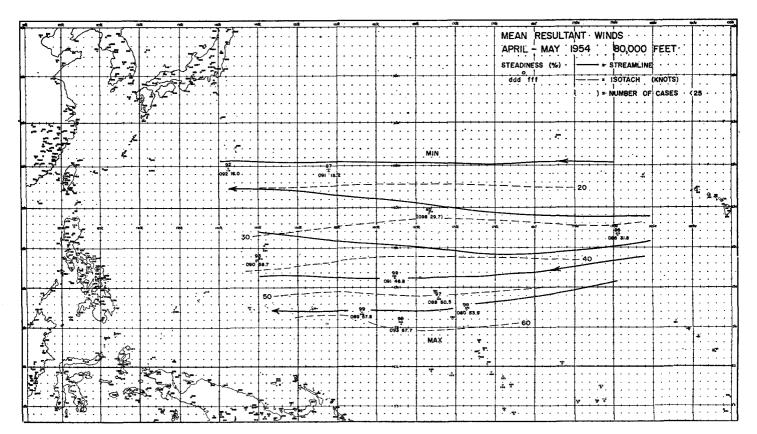


FIGURE 27.—The field of motion at 80,000 feet, April 16-May 15, 1954.

because they typify high- and low-index situations, as understood qualitatively by the synoptic meteorologist.

- 3. In the Central Pacific the wind steadiness associated with the upper tropospheric wind systems is greater than that associated with the lower tropospheric systems. This fact points up the large stability of these large-scale upper-level wind circulations in that area.
- 4. As far as the upper tropospheric and lower stratospheric circulations are concerned the maps are not incompatible with those already published (Kochanski [4], Flohn [5], U. S. Weather Bureau [6]). However, it should be pointed out that no attempt at the separation of "high-index" and "low-index" mean circulations was formerly made so that those analyses probably represent averages based on situations of both types.
- 5. There still remains the fundamental difficulty of describing in the three-dimensional picture, the relation of the circulation in the Northern Hemisphere to that of the Southern Hemisphere. This could be remedied if additional stations were instituted from 5° N. to 10° S. even for limited periods of time.

ACKNOWLEDGMENTS

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